



CALCAREOUS NANNOFOSSILS EVENTS OF NEPHROLITHUS FREQUENS ZONE TO MARKALIUS INVERSUS ZONE IN SECTION OF KAZERUN, ZAGROS (IRAN)

Saeedeh Senemari¹

¹Assistant Professor, department of Mining Engineering, Imam Khomeini International University, Qazvin, Iran

ABSTRACT

One of the most complete Late Maastrichtian to Early Danian sequences contains the Kazerun section of West Fars Province, Zagros basin of Southwestern Iran. The present investigation identifies index species, 15 Genera and 22 Species of nannofossils in this boundary. Calcareous nannofossils, which observe in Cretaceous-Paleogene boundary strata, are believed to be appropriate means for biostratigraphical study. The majority of Cretaceous-Tertiary boundary (K/Pg) strata are made up of grey shale and then purple shale, respectively. The acquired nannofossils show continuity in K/Pg boundary in Western part of Fars Province. This agrees with the existence of Nephrolithus frequens zone, which predicts the age of late Late Maastrichtian and Markalius inversus zone which describes the age of Early Paleocene (Early Danian). Also, the examination of calcareous nannofossils of the Kazerun section at Zagros basin enabled us to recognize standard biozones defined in Mediterranean regions, especially Tethysian domain.

Keywords: Nannofossil, Cretaceous, Kazerun, Paleogene, Zonation, Biostratigraphy.

Received: 12 May 2015/ **Revised:** 15 July 2015/ **Accepted:** 21 July 2015/ **Published:** 27 July 2015

Contribution/ Originality

This study is one of very few studies which have investigated based on calcareous nannofossils in Zagros. Stratification Cretaceous-Paleogene series indicate the samples in relation to Neotethys basin. Also, the environmental conditions of succession that are a part of Neotethys indicate marine and tropical environment in intervals of K/Pg at low latitudes.

1. INTRODUCTION

In order to investigate the Cretaceous-Paleogene boundary (K/Pg) by the calcareous nannofossils, the upper part of Cretaceous succession is selected in Northeast of Kazerun. This reaserch is due to the response of the calcareous nannofossils to global change across the K/Pg

that is poorly known in this area. This succession has been investigated from different aspects of paleontology in the Zagros basin (Ghasemi-Nejad *et al.*, 2006; Khosrowtehrani, 2008). Since the few studies on K/Pg boundary of calcareous nannofossils have been carried out in Southwest of Iran (Hadavi and Ezadi, 2007; Hadavi *et al.*, 2007; Senemari and Sohrabi Molla, 2012) therefore in this study we investigated the Cretaceous to Paleogene transition (K/Pg) based on calcareous nannofossils. The aim of this research is to determine the calcareous nannofossil assemblages and the possibility of using the biostratigraphic standard zonation in K/Pg, Northeast of Kazerun city.

2. GEOLOGICAL SETTING

The Zagros basin is composed of a folded rocks succession. The sediments were deposited in a gradually subsiding trough formed in the Late Triassic. It comprises of more than 10,000m of Mesozoic to Cenozoic strata (Setudehnia, 1978). The Zagros Folded belt is divided into three stratigraphic provinces that one of them is the Fars province (Motiei, 1995). The Cretaceous-Paleogene succession is exposed along the Northeast of the Kazerun of the Dashtak anticline in Fars Province (Fig. 1). The study section is located in the about 8km Northeast of Kazerun city. This section consists of 385 m thick succession of Santonian (Upper Cretaceous) together with the lower part of Paleogene strata and covers the area between longitudes $51^{\circ}40'$ - $51^{\circ}43'$ E and latitudes $29^{\circ}39'$ - $29^{\circ}41'$ N. In fact, the majority of the section consisted of Upper Cretaceous grey shale and is itself overlain by the purple shale of Pabdeh Formation. The samples were collected from the top of Upper Cretaceous and base of Pabdeh Formation.

3. MATERIAL AND METHODS

A total of 66 samples representing of the 22 m thick at interval K/Pg were obtained. The most detailed sampling was performed in the intervals at 18 m below and 4m above K/Pg boundary of the section. Samples were prepared using the smear slide technique (Young and Bown, 1998). All slides were studied under crossed-polarized and phase-contrast light microscope at $\times 1000$ magnification. The Late Cretaceous (late Late Masstrichtian) /Early Paleocene (Early Danian) boundary nannofossil species considered in this paper are referenced in Perch – Nielsen (1985a ;1985b). Abbreviations used in this study are the FO (first occurrence) and the LO (last occurrences). The NP (Nannoplankton Paleogene) zonation of Martini (1971) the CP (Coccolith Paleogene) zonation of Okada and Bukry (1980) and the CC (Coccolith Cretaceous) nannofossil zonation of Sissingh (1977) are used in the present research. The number of calcareous nannofossil species was calculated through semi quantitative analysis, with abundance code defined as follows; R = rare (1 specimen per field of view [FOV] 11-100), F = few (1 specimen per 2-10 FOV), C = common (1-10 specimens per FOV), A = abundant (>10 specimens per FOV).

4. RESULTS

In this study, stratification series of Cretaceous-Paleogene show the samples in relation to K/Pg. As a result of this study, for the first time, 15 genera and 22 species of nannofossils have

been identified. From set of slides and assemblage species, the late Late Maastrichtian (CC26 zone) to Early Danian (NP1 zone) was identified and the species ranges were determined. Marker species and the most common species are illustrated in Plate (Fig. 2). These zonal schemes are shown in Table 1. In addition, we can also learn about the environmental conditions of the succession that are in fact a part of Neotethys basin with the existence of indexed species calcareous nannofossils that indicate deep marine and tropical environment in intervals of K/Pg at low latitudes, such as the Zagros .

5. DISCUSSION

A) Nannostratigraphy and Bioevent

The calcareous nannofossils due to their wide geographic spread and fast evolutionary rate provide one of the most complete biostratigraphic zonation for Mesozoic and Cenozoic times (Bown, 1991; Burnett, 1998). The importance of these calcareous nannofossils for zonation has been discussed at length by Martini (1971); Sissingh (1977) and Perch – Nielsen (1985a ;1985b). Also these group an excellent proxy to past environmental changes from Late Cretaceous to Early Paleocene Epoch strata (Sissingh, 1977; Perch – Nielsen, 1985a; Dunkley and Bown, 2007). The most Cretaceous nannofossil taxa became extinct below the first purple shale strata, a bio event that is synchronous with the Cretaceous/Paleocene boundary (K/Pg) event in low–mid-latitude areas (Perch –Nielsen, 1985b; Lees, 2002). The evolution of various forms of calcareous nannofossil assemblages during the 1.4 million years of the late Late Maastrichtian/ equivalent with beginning of CC26 zone to Early Danian/ equivalent with beginning of NP1 zone, has enabled the recognition of two biostratigraphic zonal units on the basis of the relative ranges of index species (Perch – Nielsen, 1985a; Gradstein *et al.*, 2004). As previously mentioned, the calcareous nannofossils discussed in this research are obtained from the succession K/Pg boundary of Northeast of Kazerun. The absence of a good fundamental paleontological study based on calcareous nannofossil is the best motivation for investigating in the Kazerun section. Twenty two calcareous nannofossil species from fifteen genera were found in the course of this study. Moderately diversity and low abundance of calcareous nannofossil were found in the section and allowed recognition of the CC26 of zonation of Sissingh (1977) and NP1 of zonation of Martini (1971) or subzone CP1a of zonation of Okada and Bukry (1980). In addition, the reliability of any age determine by means of calcareous nannofossils depends on the preservation of the coccoliths (Perch –Nielsen, 1985b). In this section, the preservation of calcareous nannofossil assemblages was good. Also, the examination of calcareous nannofossils of the Northeast Kazerun section at Zagros basin enabled us to recognize standard bio zones defined in Mediterranean regions, especially Tethysian domain (Watkins, 1996). Therefore according to our biostratigraphic data, the K/Pg boundary bio zones are introduced as follows:

B) *Nephrolithus Frequens* Zone (CC26)

The first bio events recorded from grey shale of the top of Cretaceous/ late Late Maastrichtian (K/Pg) is the zone CC26 defined as the interval from the first occurrence (FO) to

last occurrences (LO) of *Nephrolithus frequens* (Perch –Nielsen, 1985b). However, in low latitudes, *Nephrolithus frequens* is not recognizable or absent but here the FO of *Micula murus* and *Micula prinsii* can be used to determine the interval between CC25 and CC26 bio zones boundary of the Upper Cretaceous (Late Maastrichtian), respectively. In other words, both species are very rare or absent in high latitudes, but are useful markers for the Late Maastrichtian in low latitudes. The assemblage in this zone is characterized by *Chiastozygus platyrhethus*, *Rhagodiscus angustus*, *Micula concava*, *Micula decussata*, *Lucianorhabdus cayeuxii*, *Lithraphidites carniolensis*, *Eiffellithus turriseiffelii*, *Eiffellithus gorkae*, *Lithraphidites quadratus*, *Micula murus*, *Micula prinsii*, *Ceratolithoides aculeus*, *Arkhangelskiella cymbiformis*, *Thoracosphaera operculata*, *Microrhabdulus decoratus*, *Braarudosphaera bigelowii*, *Arkhangelskiella maastrichtiana*, *Watznaueria biporta*, *Watznaueria barnesae*.

C) Markalius Inversus Zone (NP1)

The other unit recorded in the formation is the *Markalius inversus* zone that is synchronous with K/Pg boundary in low latitude areas. This zone is defined as the interval from the last occurrences (LO) of Cretaceous coccoliths or FO of acme of *Thoracosphaera* (*T. operculata*) bloom to FO of *Cruciplacolithus primus*. These species have been recorded slightly above the K/Pg boundary in the Early Paleocene (Early Danian). The NP1 assemblage is characterized by *Biantholithus sparsus*, *Markalius inversus*, *Cruciplacolithus primus*, *Arkhangelskiella cymbiformis*, *Micula decussata*, *Thoracosphaera operculata*, *Micula prinsii*, *Braarudosphaera bigelowii*.

6. CONCLUSION

The Plate illustrates the calcareous nannofossils diversity data for the K/Pg boundary in the Northeast of Kazerun city. The species include: *Eiffellithus turriseiffelii*, *Micula decussata*, *Cruciplacolithus primus*, *Biantholithus sparsus*, *Watznaueria barnesae*, *Watznaueria biporta* are the major components and abundance. *Ceratolithoides aculeus*, *Micula murus*, *Arkhangelskiella cymbiformis*, *Arkhangelskiella maastrichtiana*, *Thoracosphaera operculata*, *Micula prinsii* and *Braarudosphaera bigelowii* are relatively abundant. *Lucianorhabdus cayeuxii*, *Chiastozygus platyrhethus*, *Rhagodiscus angustus*, *Micula concava*, *Microrhabdulus decorates*, *Markalius inversus*, *Lithraphidites carniolensis*, *Lithraphidites quadratus* are rare. Based on the nannofossils are listed, age the studied upper section is late Late Maastrichtian that corresponding to CC26 Zones (Sissingh, 1977) to zone NP1 of zonation of Martini (1971). Species *Biantholithus sparsus*, *Markalius inversus*, *Cruciplacolithus primus*, *Micula prinsii*, *Braarudosphaera bigelowii* obtained in the upper cretaceous layers, in across the Cretaceous-Tertiary boundary, which to explain beginning of Paleocene.

Funding: This study received no specific financial support.

Competing Interests: The author declares that there are no conflicts of interests regarding the publication of this paper.

REFERENCES

Bown, P.R., 1991. Calcareous nannofossil biostratigraphy. London: Chapman and Hall.

- Burnett, J.A., 1998. Upper cretaceous calcareous nannofossil biostratigraphy. In: Bown, P.R. (Eds). London: Chapman & Hall/Kluwer Academic. pp: 199-132.
- Dunkley, J.T. and P.R. Bown, 2007. Post-sampling dissolution and the consistency of nannofossil diversity measures: A case study from freshly cored sediments of coastal Tanzania. *Marine Micropaleontology*, 62(4): 254–268.
- Ghasemi-Nejad, E., M.H. Hobbi and P. Schioler, 2006. Dinoflagellate and foraminiferal biostratigraphy of the gurpi formation (Upper Santonian-Upper Maastrichtian). *Zagros Mountains, Iran. Cretac Res*, 27(6): 828–835.
- Gradstein, F.M., J.G. Ogg, A.G. Smith, F.P. Agterberg, W. Bleeker, R.A. Cooper, V. Davydov, P. Gibbard, L.A. Hinnov, M.R. House (†), L. Lourens, H.P. Luterbacher, J. McArthur, M.J. Melchin, L.J. Robb, P.M. Sadler, J. Shergold, M. Villeneuve, B.R. Wardlaw, J. Ali, H. Brinkhuis, F.J. Hilgen, J. Hooker, R.J. Howarth, A.H. Knoll, J. Laskar, S. Monechi, J. Powell, K.A. Plumb, I. Raffi, U. Röhl, A. Sanfilippo, B. Schmitz, N.J. Shackleton, G.A. Shields, H. Strauss, J. Van Dam, J. Veizer, T. Van Kolfshoten and D. Wilson, 2004. A geologic time scale 2004. Cambridge, United Kingdom (UK): Cambridge University Press.
- Hadavi, F. and M. Ezadi, 2007. Biostratigraphy of the gurpi formation in dare-shahr section (Zagros Basin). The First MAPG International Convention Conference and Exhibition. pp: 28–31.
- Hadavi, F., K. Khosrowtehrani and S. Senemari, 2007. Biostratigraphy of calcareous nannofossils of gurpi formation in North Gachsaran. *J Geosci*, 64(16): 14–23.
- Khosrowtehrani, K., 2008. Applied micropaleontology. Iran: Tehran University Press.
- Lees, J.A., 2002. Calcareous nannofossil biogeography illustrates palaeoclimate change in the late cretaceous Indian ocean. *Cretaceous Research*, 23(5): 537–634.
- Martini, E., 1971. Standard tertiary and quaternary calcareous nannoplankton zonation. In: Farniacci, A. (Eds). *Planktonic Microfossils. Second Planktonic Conference Proceedings*, Edizioni Tecnoscienza, Rome. pp: 739–785.
- Motiei, H., 1995. Petroleum geology of zagros-1. Geological survey of Iran with cooperation of deputy ministry of project and planning, Tehran, Iran (In Persian), No. 25. Available from http://www.petroshop.ir/product.php?id_product=185&
- Okada, H. and D. Bukry, 1980. Supplementary modification and introduction of code numbers to the low-latitude coccolith biostratigraphic zonation. *Marine Micropaleontology*, 5(c): 321–325.
- Perch – Nielsen, K., 1985a. Cenozoic calcareous nannofossils. In: Bolli, H. M., Saunders, J. B., Perch-Nielsen, K. (Eds). *Plankton stratigraphy*. Cambridge: Cambridge University Press, 1: 427–554.
- Perch –Nielsen, K., 1985b. Mesozoic calcareous nannofossils. In: Bolli, H. M., Saunders, J. B., Perch-Nielsen, K. (Eds). *Plankton stratigraphy*. Cambridge: Cambridge University Press, 1: 329- 426.
- Senemari, S. and U.M. Sohrabi Molla, 2012. Evaluation of cretaceous–paleogene boundary based on calcareous nannofossils in section of Pol Dokhtar, Lorestan, South Western Iran. *Arabian Journal of Geosciences*, 6(10): 3615–3621.
- Setudehnia, A., 1978. The mesozoic sequence in South-West Iran and adjacent areas. *Journal of Petroleum Geology*, 1(1): 3-42.

- Sissingh, W., 1977. Biostratigraphy of cretaceous calcareous nannoplankton. *Geologie En Mijnbouw*, 56(1): 37-65.
- Watkins, D.K., 1996. Upper cretaceous calcareous nannofossil biostratigraphy and paleoecology of the Southern ocean. In: Alicia, M., Robin, W. (Eds). *Microfossils and oceanic environments*. Aberystwyth: University of Wales Press. pp: 355-381.
- Young, J.R. and P.R. Bown, 1998. Cenozoic calcareous nannoplankton classification. *Journal of Nannoplankton Research*, 19(1): 36-47.

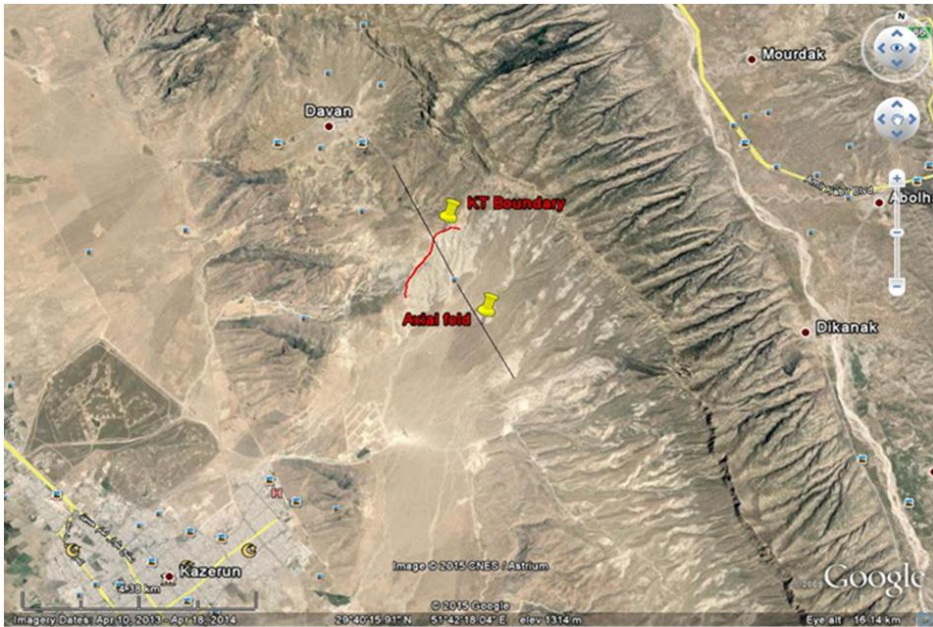


Figure-1. Geographical location of anticlinal, after Google (2014)

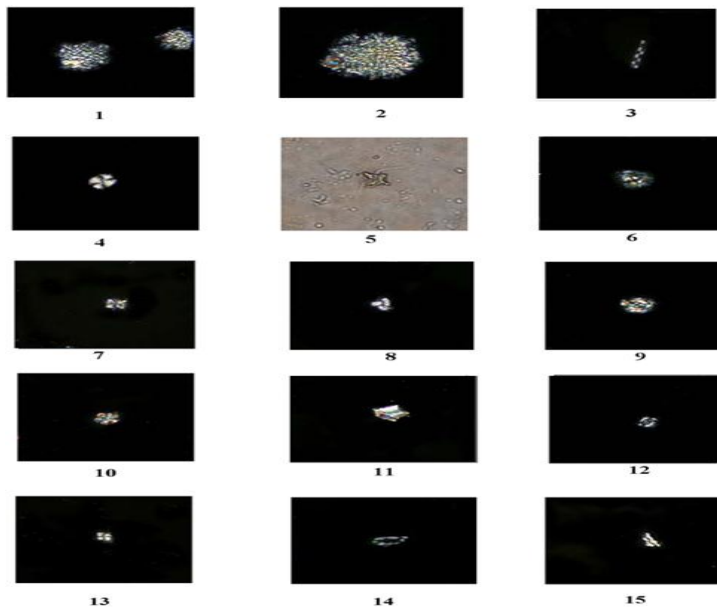
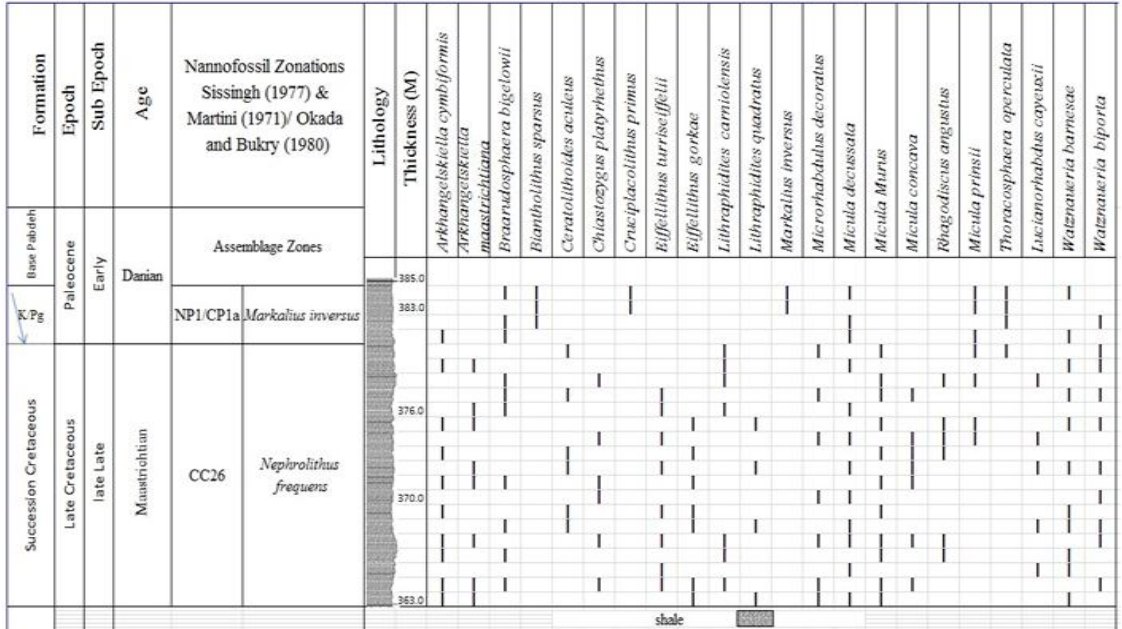


Figure-2.Plate: All figures in XPL except fig.5 in PPL, light micrographs×1000; the taxa considered in the present Figure is referenced in (Perch – Nielsen, 1985a ;1985b). 1, 2: *Thoracosphaera operculata*; 3: *Microrhabdulus decoratus*; 4: *Watznaueria biporta*; 5: *Micula concava* (Stradner in; 6: *Markalius inversus*; 7: *Micula decussata*; 8: *Miculaprinisii*; 9: *Cruciplacolithus primus*; 10: *Biantholithus sparsus*; 11: *Braarudosphaera bigelowii*; 12: *Cruciplacolithus primus*; 13: *Micula murus*; 14: *Rhagodiscus angustus*; 15: *Lucianorhabdus cayeuxii*.

Table-1. The chart of Nanno stratigraphic K/Pg boundary (Northeast of Kazerun)



Views and opinions expressed in this article are the views and opinions of the author(s), International Journal of Geography and Geology shall not be responsible or answerable for any loss, damage or liability etc. caused in relation to/arising out of the use of the content.